TECGREMED



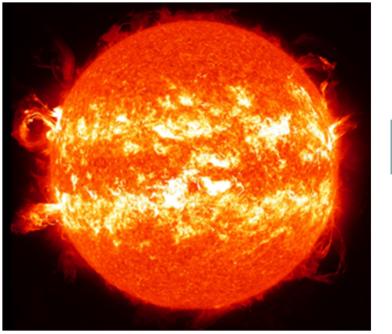
València, Spain | 26–28 may2025

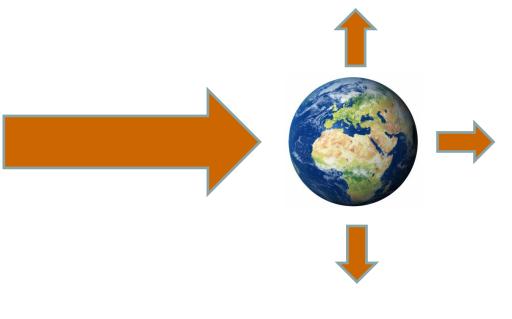
Properties of Solar Energy

Instructor: Miguel Ángel Pérez Martín, e-mail: <u>mperezm@hma.upv.es</u> DIHMA room nº 22



1. Solar properties



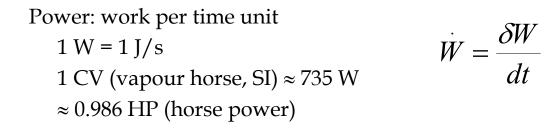


Basic units on Energy



Energy (J) = Work done (Force x distance, Nm) = Power x time (W x s)

Work = Energy: $\delta W = \vec{F} \cdot d\vec{S} = p \cdot dV$ 1 J (Julio) = 1 Nm (Newton x metro) = 1 Ws (Watt x second) = 0.2388 cal ≈ 0.24 cal1 kWh = 860 kcal = 3,600 kJ1 tep (tone oil equivalent toe) = 11,63 MWh



Installed power: the capacity to produce energy, the measurement unit is W (kW, MW, GW) **Energy production:** kWh, MWh, GWh

light bulb 100 W (discontinued) consume => per year (8760 h/year) 876 kWh => 876kWh-year ≈ 1,000 kWh-year = 1MWh-año Power(production capacity)kW kilowatt = 10^3 WMW Megawatt = 10^6 WGW Gigawatt = 10^9 WTW Terawatt = 10^{12} WPW Pentawatt = 10^{15} WEW Exawatt = 10^{18} W

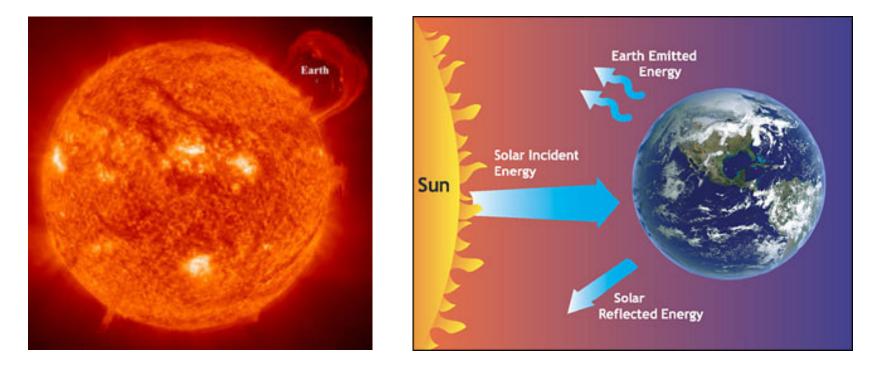




Energy from the Sun



Energy inputs: Solar Radiation – Distance -Atmosphere



Energy depends on: Emission Law: Stefan-Boltzmann Law Energy distribution: Planck Law Sun Radiation



Sun Photosphere temperature 5.785 °K (5.800), Sun Black body => => electromagnetic radiation Stefan-Boltzmann Law. Josef Stefan (1835 - 1893) y Ludwig Boltzmann (1844 - 1906)

 $E = \varepsilon \cdot \sigma \cdot T^4$

σ = 5.669 x 10⁻⁸ Wm⁻² °K⁻⁴ Stefan-Boltzmann constant ε emissivity (black body, ε = 1, general ε ≤ 1)

Sun surface emission assuming black body: $E = 6.35 \times 10^7 Wm^{-2}$

Total Sun Emission. Sun radius, r_s , 6.959 x 10⁸ m, As= $4\pi r_s^2$ $E_s = E * A_s = 3.873 \times 10^{26} W$ Solar Irradiance: energy received from Sun

Total Sun Emission.

$$L_s = 4 \cdot \pi \cdot r_s^2 \cdot \sigma \cdot T^4 = 3.65 \cdot 10^{26} W$$

Solar emission sphere $S = 4 \cdot \pi \cdot a_0^2$

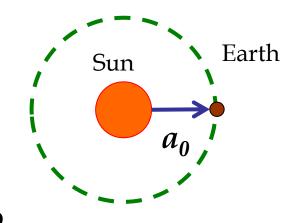
Irradiance: solar radiation in the top of the atmosphere

$$I_{SC} = S_0 = \frac{L_s}{S} = \sigma \cdot T_S^4 \cdot \left(\frac{r_s}{a_0}\right)^2 = 1366 \frac{W}{m^2} \quad \text{Solar constant}$$

Ts = 5.776 K sun temperature

 r_s Sun radius: 6.959 x 10⁸ m

 a_0 Sun-Earth distance (1 UA, 149.597.871 km, irradiance arrives in 8 min y 19 s) σ Stefan-Boltzmann constant (4.903 10⁻⁹ MJ m⁻² K⁻⁴ day⁻¹)

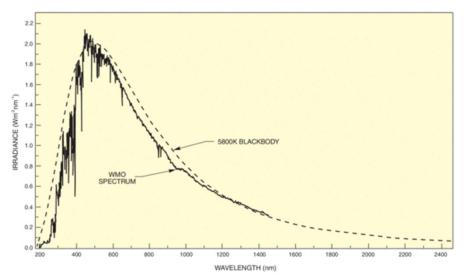




Spectral distribution wavelenght, λ, Planck's Law German physic Max Planck (1858 - 1947)

$$c_1 = 2\pi h c^2 \qquad c_2 = h c / k_{\rm B}.$$

first radiation constant c_{1L} and the second radiation constant c_2 h Planck constant => h = 6,62606957(29) ×10 -34 J×s



Speed of light c = 299,792,458 m/s

Boltzmann's constant relates energy and temperature k = $1.3806504 \times 10^{-23}$ J/K

Electro magnetic spectrum

Top of atmosphere irradiance (Iqbal, 1983) $W/m^2/\mu m^1$ Dotted line: from Plank Law assuming Sun temperature 5785°K Continuous line: observed irradiance from satellites and rockets

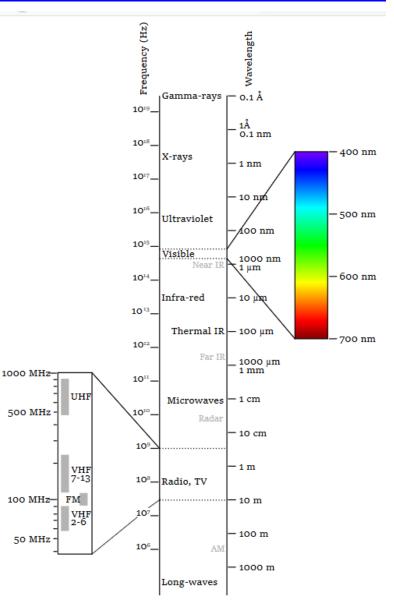


 $E_{\lambda} = \frac{1}{\lambda^5 \cdot \left(e^{\frac{C_2}{\lambda T}} - 1\right)}$



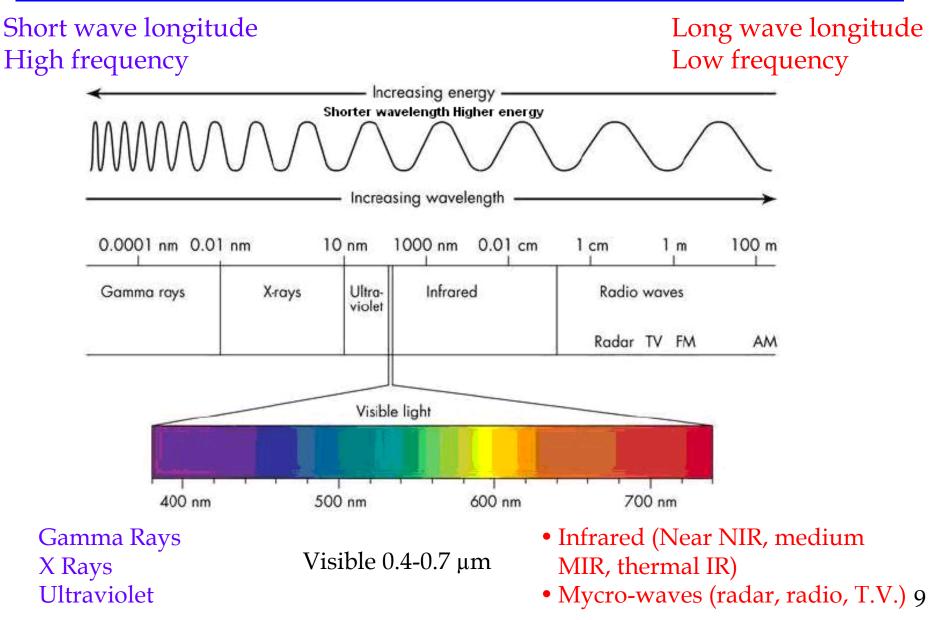
Electromagnetic spectrum

- Ionizing radiation:
 - ➤ Gamma rays
 - Ultraviolet: short wave
- Visible: sensible to human eye.
 Three main areas: blue, green and red. (0.4 0.78 μm)
- > Infrared:
 - ➢ NIR Near infrared 1 µm
 - ➢ MIR Medium Infrared 10 µm
- Micro-waves and radio waves



Electromagnetic spectrum

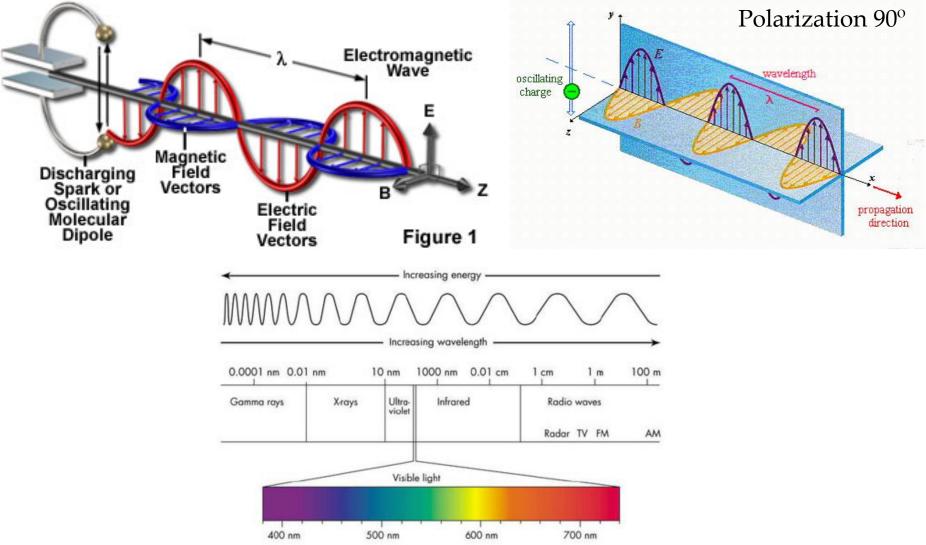




Electromagnetic waves

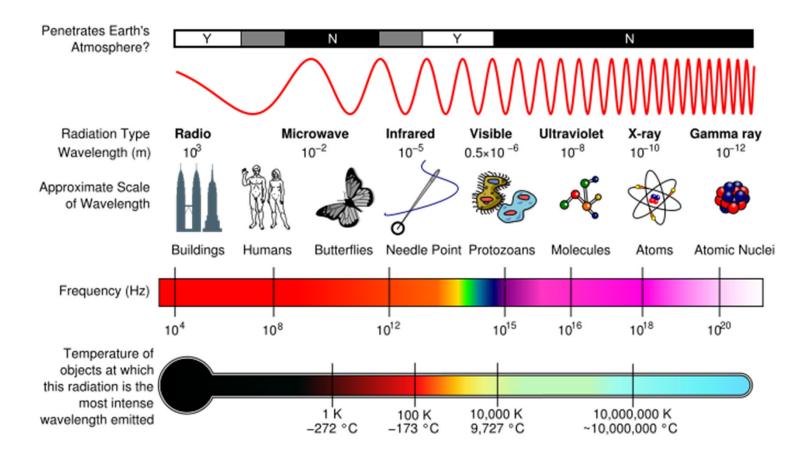


Propagation of an Electromagnetic Wave



Electromagnetic waves properties





Solar energy and wavelength

2500



Energy in the top of atmosphere TOA

TOA: a given altitude where air becomes so thin that atmospheric pressure or mass becomes negligible

Spectral Irradiance W*m-2000 2*micron-1 1500 1000 500 0 0.00 0.501.001.503.502.00 2.50 3.00 4.00Wavelength micrometers • E-490 W/m2/micron 1600 Spectral Irradiance W*m-1400 1200 2*micron-1 1000 800 600 400 200 0 0.00 0.501.001.502.50 3.50 4.002.00 3.00 Wavelength micrometers • E-490 W/m2/micron

Solar energy in the top of atmosphere 1.366 – 1.367 Wm⁻²

Solar constant

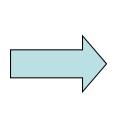
Integrating for all the wavelengths

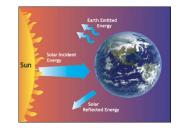
Energy from the Sun

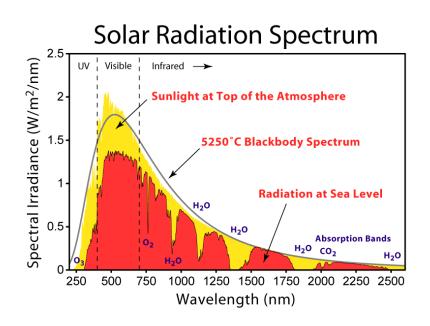


Irradiance in the TOA (Top of Atmosphere): 1.367 W/m^2 Irradiance at sea level with clear sky 1.000 W/m^2

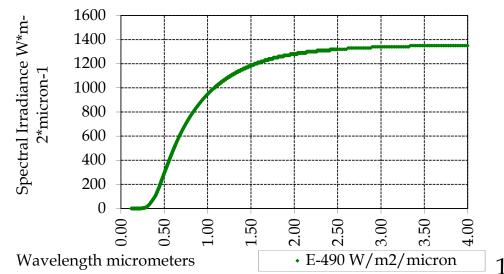








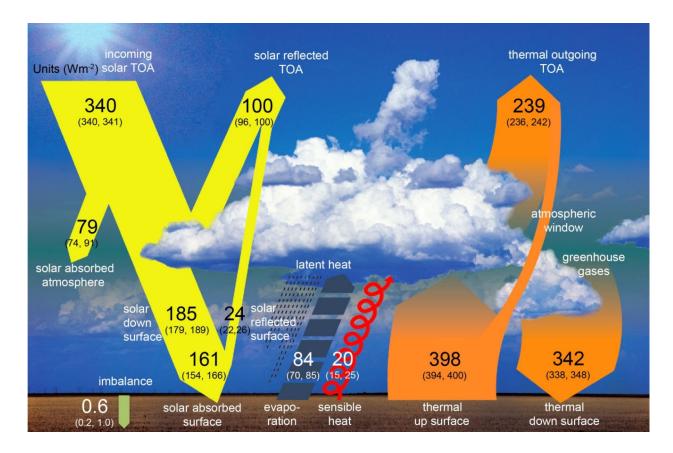
Integrated irradiance from Sun



13



Atmosphere interaction





Reflectivity: 35% of incoming solar radiation is reflected back to space).

- Clouds; Twenty-four percent of incoming solar radiation is reflected by clouds, 4% by the Earth's surface
- Scattering; Seven percent of incoming solar radiation is scattered back to space. Particles in the atmosphere can scatter incoming solar radiation

Absorption: about 17% of incoming solar radiation is absorbed at various levels in the atmosphere. Absorption is the process by which radiant energy is transferred to matter

Gases, therefore, are not like black bodies that absorb equally and completely at all wavelengths. Rather, they absorb only at specific, often narrow ranges of wavelengths. Diatomic molecules such as nitrogen and oxygen (most of our atmosphere) can absorb energy by increasing the vibration of the bond between the two atoms. If the energy absorbed is great enough it may break the bond resulting in two free wheeling oxygen or nitrogen atoms traveling at high speeds.

Ozone absorption

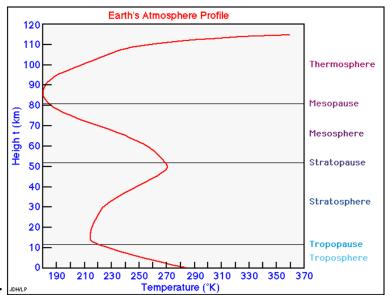


- In the highest regions of the atmosphere 110-100 km Oxygen molecules are therefore broken into oxygen atoms => O2 + ultraviolet light = O + O
- For some distance above and below 80 kilometers there is little absorption of solar energy and consequently little heating of the atmosphere so the temperature reaches a minimum.
- Descending below eighty kilometers the atmosphere is heated by another process. The atmosphere gets denser (thicker) and with decreasing altitude form an ozone molecule => O + O2 = O3
- Ozone can in turn be broken up by ultraviolet light, 50-10 km, resulting in this reaction => O3 + ultraviolet light = O2 + O
 Earth's Atmosphere Profile

Both:

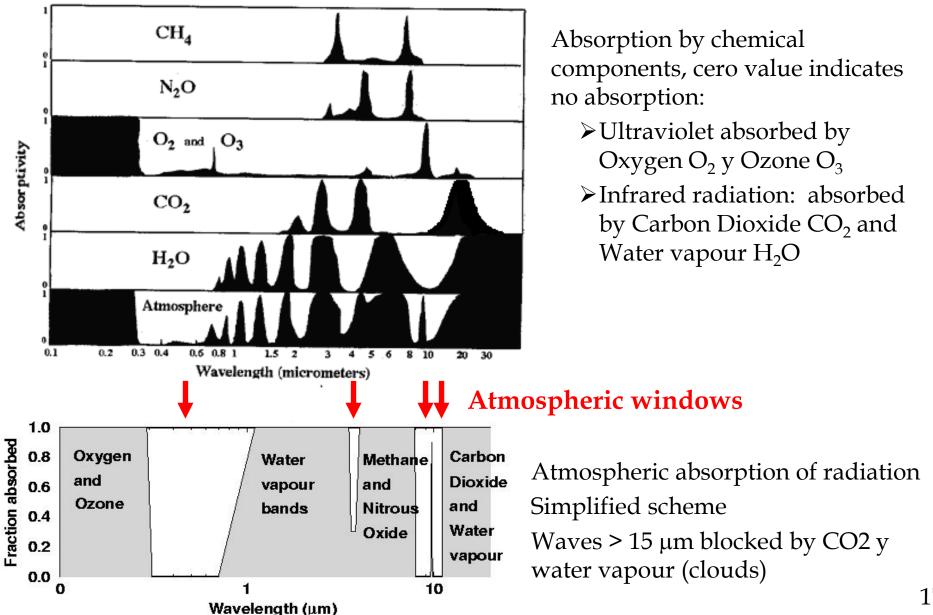
- the breaking up of oxygen molecules above fifty kilometers
- and ozone molecules at fifty kilometers and below.
 Causes heating of the atmosphere that peaks at about 50 kilometers (the stratopause).

Between 50 and 10-15 kilometers (the stratosphere) the solar energy energetic enough to break up ozone (ultraviolet radiation) is used up and the atmosphere cools.

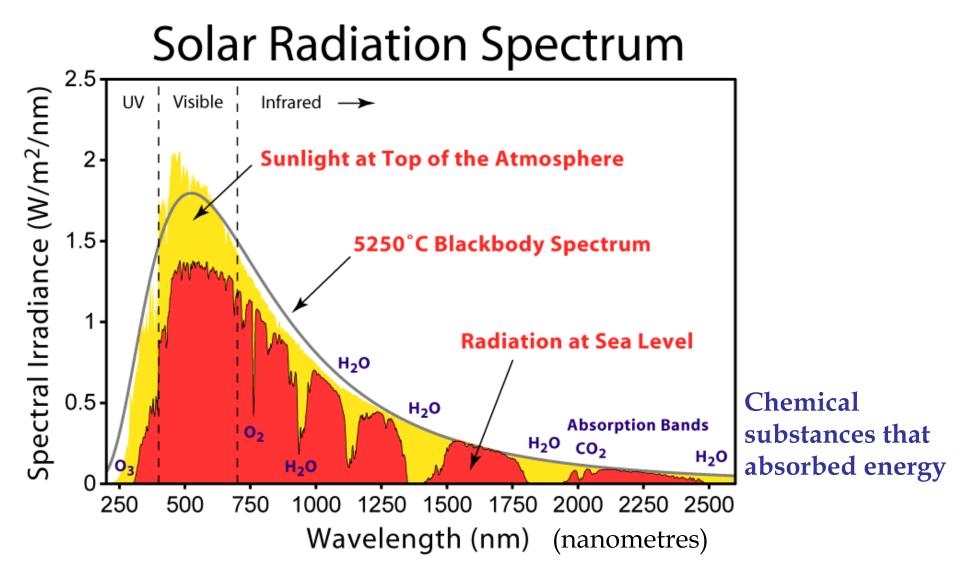


Energy absorption by chemical components

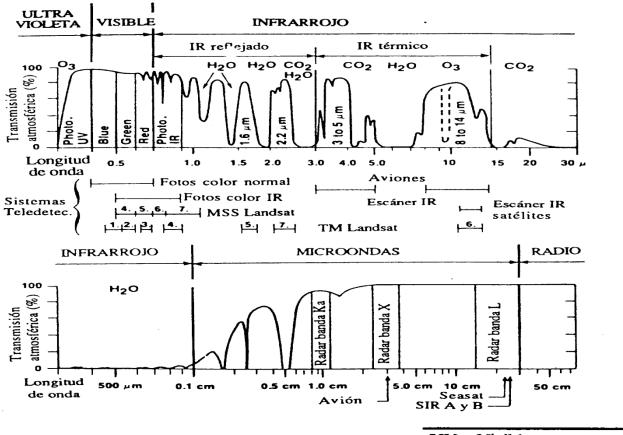








Energy Transmittance in the atmosphere and atmospheric windows



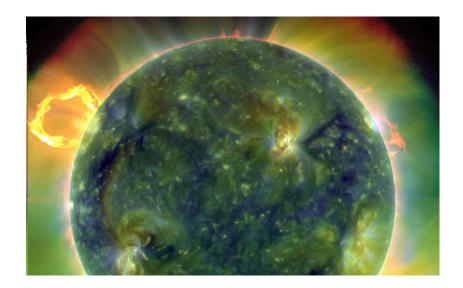
UV y Visible	0,30 - 0,75 μm
IR cercano	0,77 - 0,91 μm
	1,55 - 1,75 μm
	2,05 - 2,4 μm
IR térmico	8,0 - 9,2 μm
	10,2 - 12,4 μm
Microondas	7,5 - 11,5 mm
	> 20 mm



Solar variation

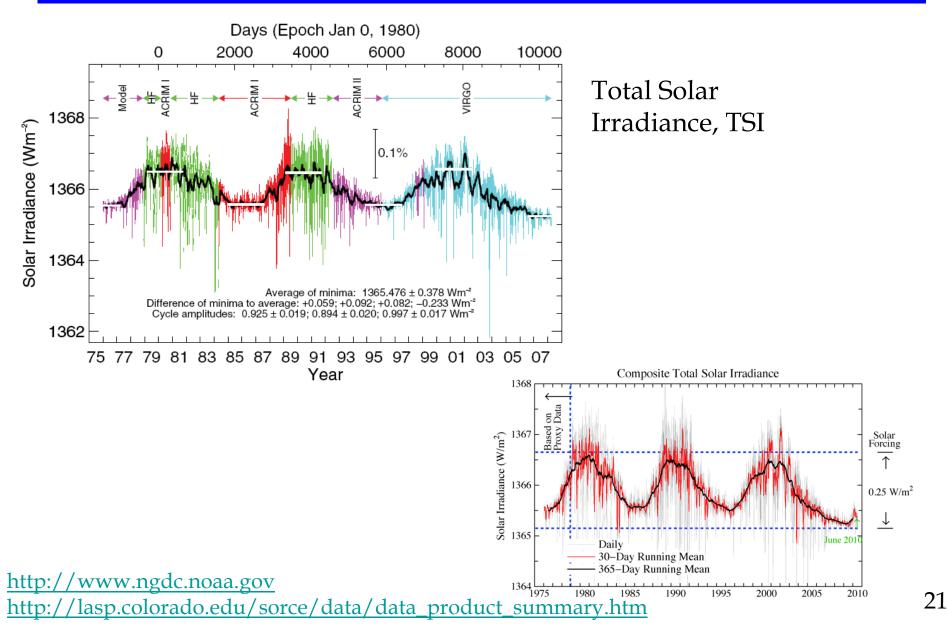






Satellite SOHO. ESA-NASA Solar and Heliosphere Observatory Launched 2-Dic-1995

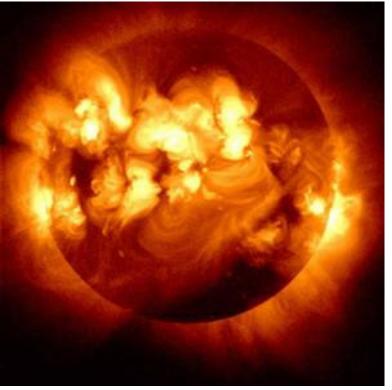




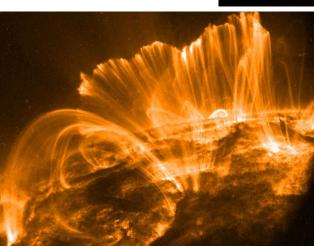
Variation in the Solar activity. Solar cycles



The Solar activity has variations and the Solar Irradiance that reach the Earth also has these variations









Sunspots are temporary phenomena on the Sun's photosphere that appear as spots darker than the surrounding areas

"Sunspots" historical records



"Sunspot number" computed daily

the sunspot number (Rudolph Wolf in 1848): R = (10*G + S)*K

G = the number of sunspot groups observed

S = the count of all sunspots in all groups

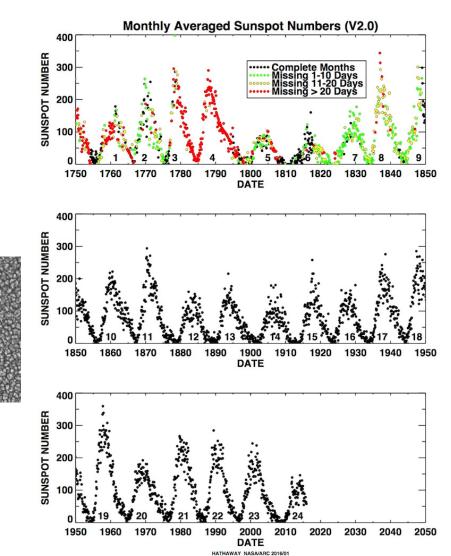
K = a scaling number to compensate for variables To combine different observers

11 years cycles

Wolf number (1864) (International Sunspot Number or Zurich number)

Actually (from 1990), only groups are counted. Two International Sunspot Number "official":

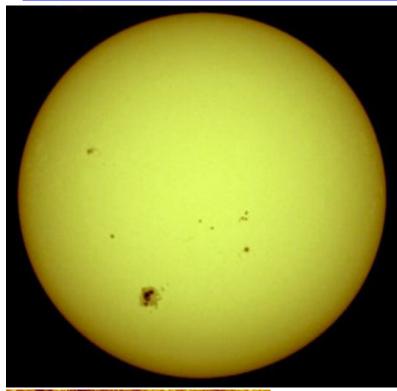
- by Sunspot Index Data Centre in Belgium.
- NOAA by US National Oceanic and Atmospheric Administration.



https://www.astroleague.org/content/what%E2%80%99s-sunspot-number

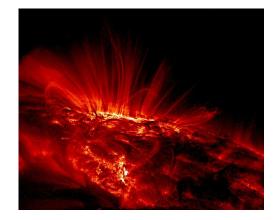
Sunspots and Solar Radiation

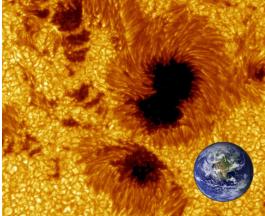




Temperature in the photosphere 5700 K Temperature in the centrum of the darkness areas (Sunspots) 3700 K

Sunspots may last a few days, and larger ones may last a few weeks.





Granules are relatively small (around 1000 km)

Relative size of the Earth with a sunspot

Sunspots and climate. Maunder minimum 1650-1700

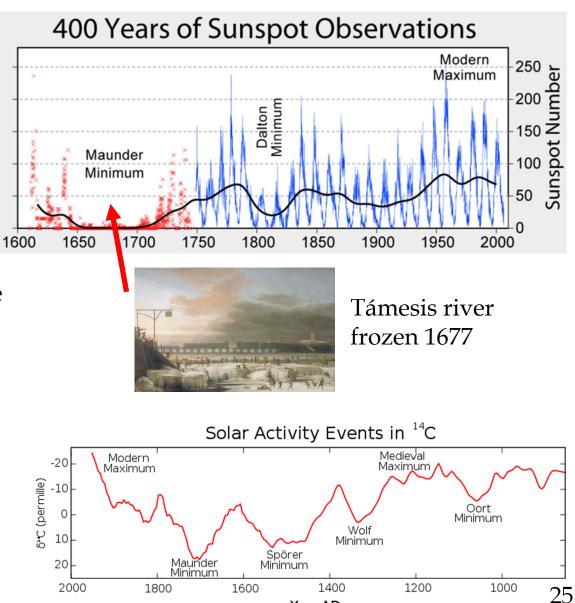


Maunder minimum. Low Solar activity between 1645 y 1715, "Little Ice Age"

Evidence of the relationship between solar activity and climate on earth. Minimum of Dalton (1790-1820)

year without summer 1816 (ash from Tambora volcano).

Reconstruction of the existence of other previous solar minima from carbon 14 in the rings of trees. Minimum of Spörer (1450-1540)



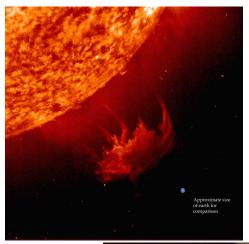
Year AD

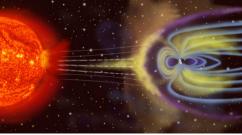


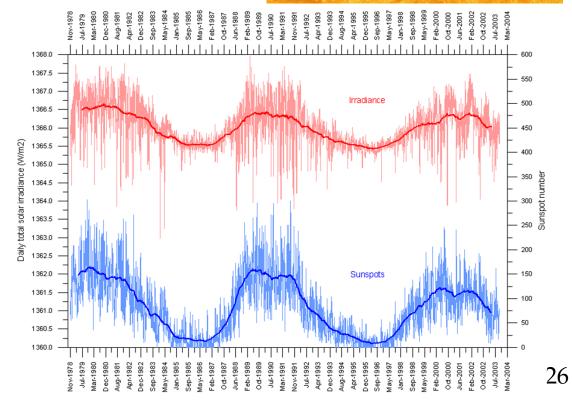


- The greater number of sunspots increases the energy that reaches the earth. Sunspot is due to a solar explosion, around the sunspot is brighter and emits more energy
- The sunspot is a relatively cold zone but due to a large heat output (flare)
- Solar explosion or solar storm occur around sunspots





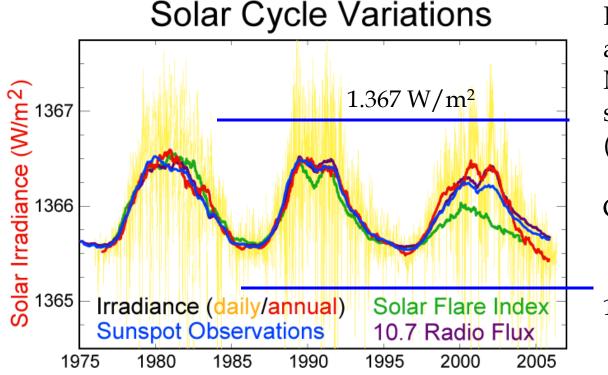




Solar Energy reaches the top of atmosphere

Solar constant at one AU, reach the Earth (Commission for Instrumentation and Observation Methods, October de 1981) Oscillation between 1.365 y 1.367 Wm⁻². Consensus 1367 ± 4 Wm⁻²

$$I_{SC} = 1367Wm^{-2} = 4.921MJm^{-2}h^{-1}$$



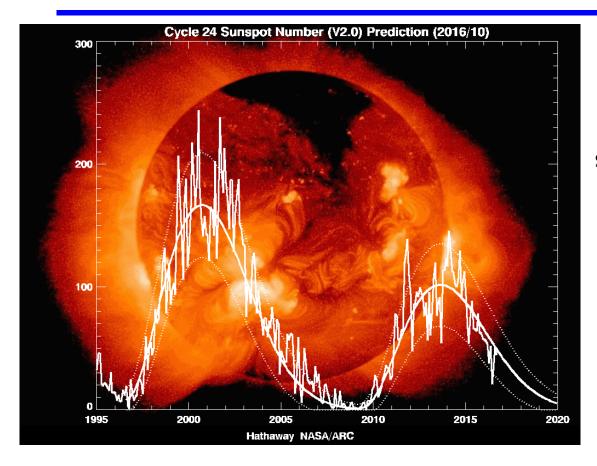
Now is 1361 W/m2

Energy variation at top of atmosphere 2 W/m² Mean around the entire surface of Earth $(1/4) => 0.5 W/m^2$

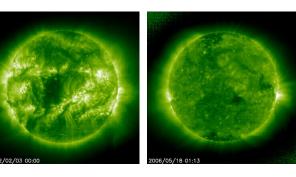
Greenhouse gases GHG Actual **1.5 W/m²** 2100: **4-8 W/m²** 1.365 W/m²

Future evolution of Sunspots



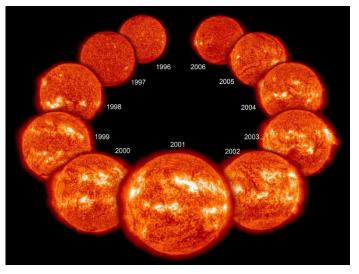


Maximum 2000 a 2002



Minimum 2006 - 2010 Cycle nº 24

Rayos-x Solar Radiation, active phase (1991, left) until inactive phase (1995, right) Satellite images from Yohkoh each 4 months. Satellite catastrophic failure 2001 (NASA, Yohkoh Science Project, <u>YPOP</u>)



<u>http://solarscience.msfc.nasa.gov/</u> <u>http://solarscience.msfc.nasa.gov/</u> <u>/SunspotCycle.shtml</u> 28

Solar constant



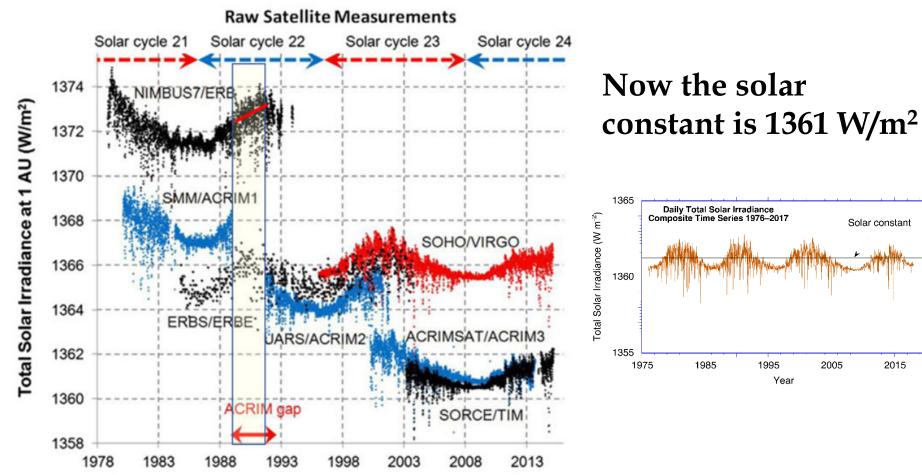


Figure 2. Raw satellite total solar irradiance (TSI) measurements. The ACRIM gap is identified in yellow. The trend of the NIMBUS7/ERB instrument in the ACRIM gap is emphasized with a red line. Source: (Soon, Connolly and Connolly 2015).

More precise measurements ACRIM (Active Cavity Radiometer Irradiance Monitor)



2. Potential Solar Energy





Energy from the Sun: Irradiance

Irradiance (power) in a surface, depends on:

- Solar angle (day of the year, hour of the day)
- ≻Weathers (cloudiness)
- ➤Surface orientation

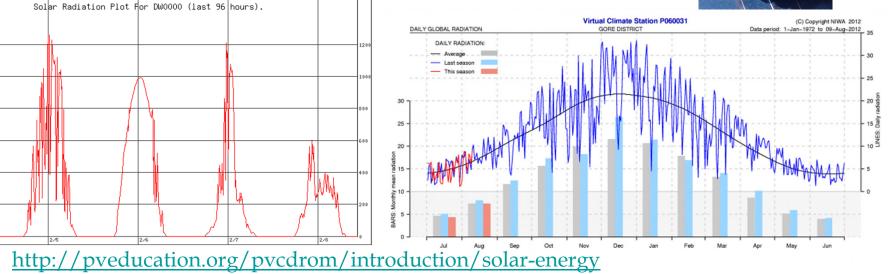
Irradiance data (hourly and daily)

- From direct observations (pyranometer)
- From theatrical estimations: latitude, cloudiness











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Energy from the Sun: Irradiance



A **pyranometer**, used to measure global irradiance



A **pyrheliometer**, mounted on a solar tracker, is used to measure Direct Normal Irradiance (or beam irradiance)



Direct Normal Irradiance (DNI)

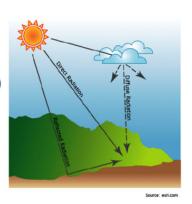


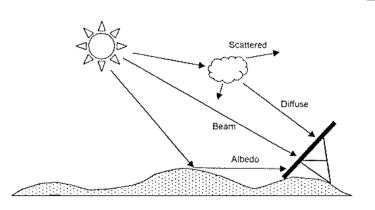
• Diffuse Horizontal Irradiance (DHI)



• Global Horizontal Irradiance (GHI)







Daily Irradiation components kWh

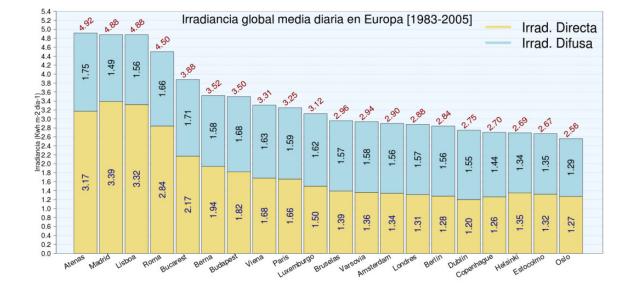
Irradiation (kWh): the sum of the total irradiance (kw) in a day

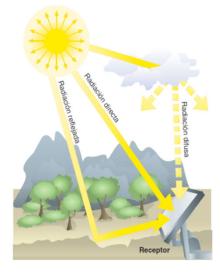
"Global daily Irradiation" Gd the sum of:

- <u>Direct radiation</u> Id:
- <u>Diffuse Radiation</u> Dd: disperseb by the atmosphere and clouds.
- <u>Reflected Radiation "albedo"</u>: reflected by the land surface. (It is usually negligible or is integrated into the diffuse)

Irradiation in Europe: The diffuse is similar

1.5 kWh/m²-día Direct changes more Spain 3.5 kWh/m²-día





Diffuse irradiation Estimation in a horizontal surface D_d (KWh/m² día) $D_d = G_d [1.39-4.027K_d + 5.531 K_d^2 - 3.108 K_d^3]$ G_d = global irradiation observed in a horizontal surface K_d = Atmospheric transparency index = G_d/G_{0d} G_{0d} = Solar irradiation at the TOA in a horizontal surface



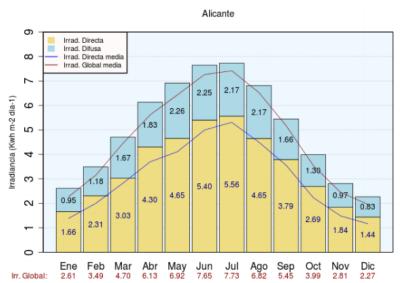
 $G_d = I_d + D_d$

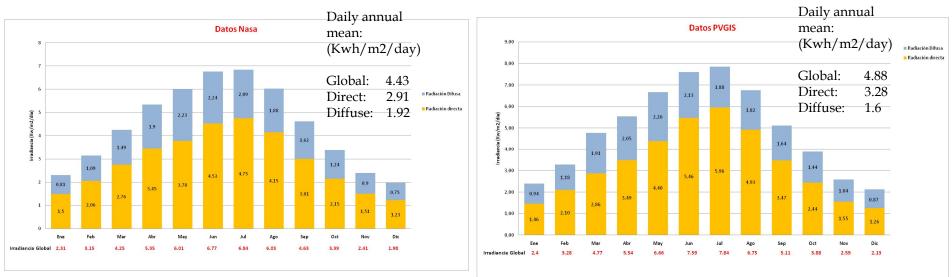
Diffuse and direct Irradiation (kWh)



Mean Irradiation in a normal day

It is lower than the irradiation day clear, because it is considered a cloudy percentage of days





Example: Alicante data

Energy from the Sun



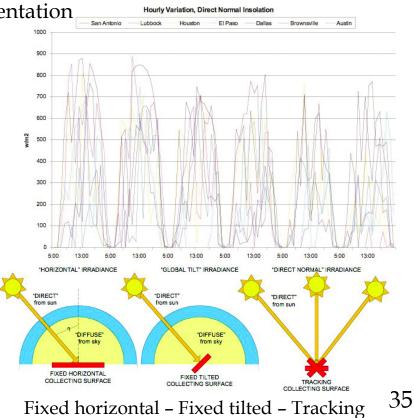
Potential Solar Energy: 3,9×10⁶ EJ/year, global irradiance over the earth (land and seas). 100% surface occupied and 100% efficiency of the solar modules.

- ▶ Realistic values: only the land and real efficiencies: 1.575 49.837 EJ/year
- ➤3 to 100 times the global energy demand on the Earth. Earth's demand in 2008, 500 EJ, in 2050 1.000 EJ.

Irradiation and electrical production depends on: In Latitude (hours of sun) – Cloudiness- Solar module orientation

Irradiance (Power) kW/m2

W/m ²	Solar constant TOA	1,361
W/m ²	Irradiance at sea level in a clear day	1,000
kW/m ²	Irradiance in the middle day	1
h/day	h/day equivalent: Valencia (variable during the day, latitude, cloudiness)	5.5
kWh/m²-día	Irradiation	5.5
días	Días/año	365
kWh/m²-año	Irradiation (annual)	2,008
	performance	0.75
kWh/m²-año	Potential annual production	1,506

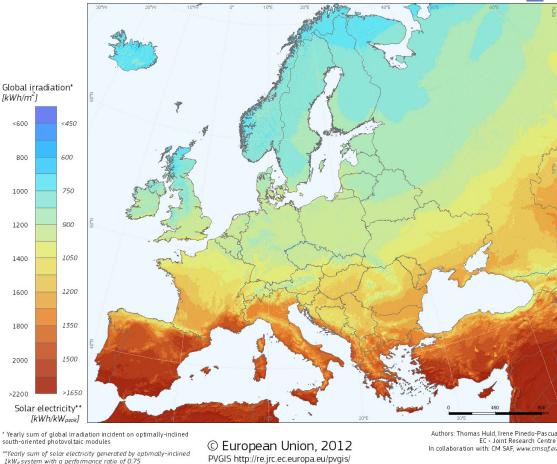


Annual photovoltaic potential in Europe kWh/m²-año Irradiation



Irradiation (Insolation): Total energy in a day kWh/m2/day

Photovoltaic Solar Electricity Potential in European Countries



Legal notice: Neither the European Commission nor any person acting on behalf of the Commission is responsible for the use which might be made of this publication Annual production depends on:

- distribution of sunshine hours during the year
- and the irradiance at those hours

Spain: (Fixed tilted)

kWh/m ² - year	min	max
Irradiation	1,400	2,000
Electricity potential production (x 0.75)	1,050	1,500

Šúri M., Huld T.A., Dunlop E.D. Ossenbrink H.A., 2007. Potential of solar electricity generation in the European Union member states and candidate countries. Solar Energy, 81, 1295–1305, http://re.jrc.ec.europa.eu/pvgis/.

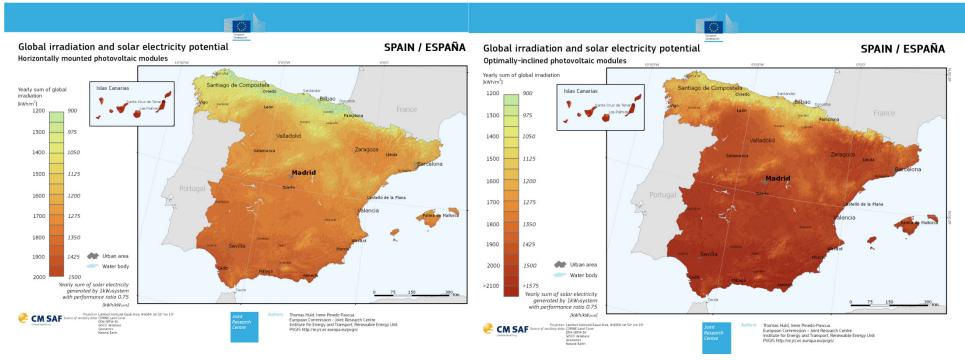
Huld T., Müller R., Gambardella A., 2012. A new solar radiation database for estimating PV performance in Europe and Africa. Solar Energy, 86, 1803-1815. 36

Annual value of horizontal irradiation and with the optimum inclination 1998-2011. kWh/m²



Fixed horizontal

Fixed tilted

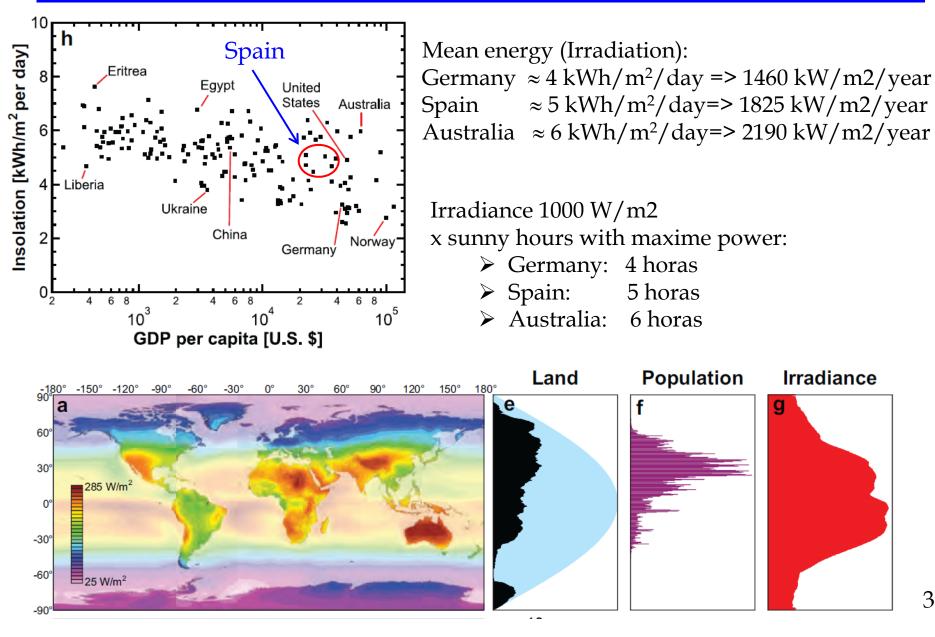


Potential Solar Electricity) [kWh/kWp] generated by a installed power of 1 kWp. Assuming System performance 0.75.

Potential electricity \approx 1,400 kWh/m² Potential annual production 1% surface (5,000 km²) 7,000 TWh Energetic demand in Spain 1,500 TWh (5 times the energetic demand)

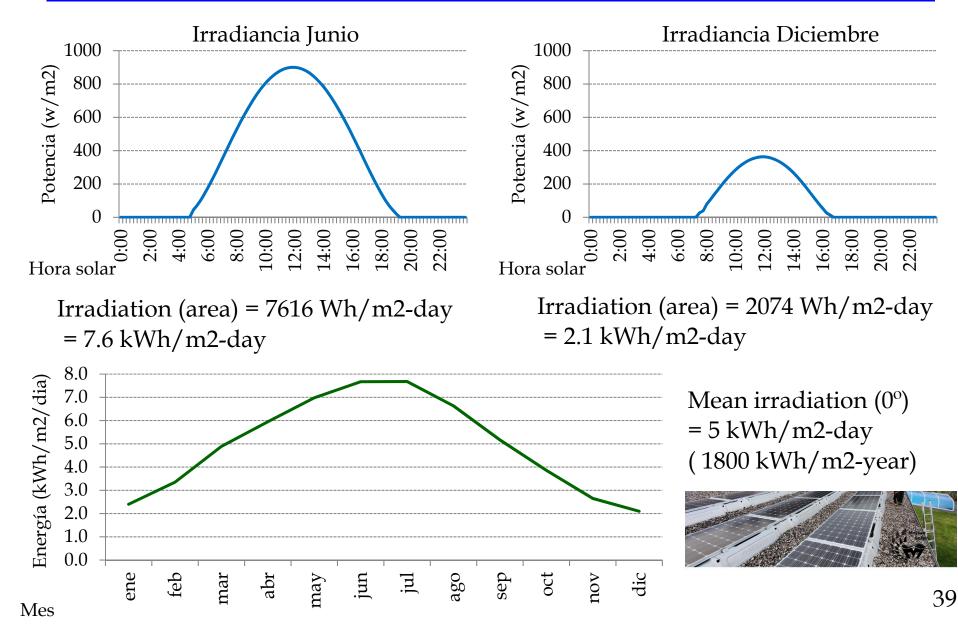


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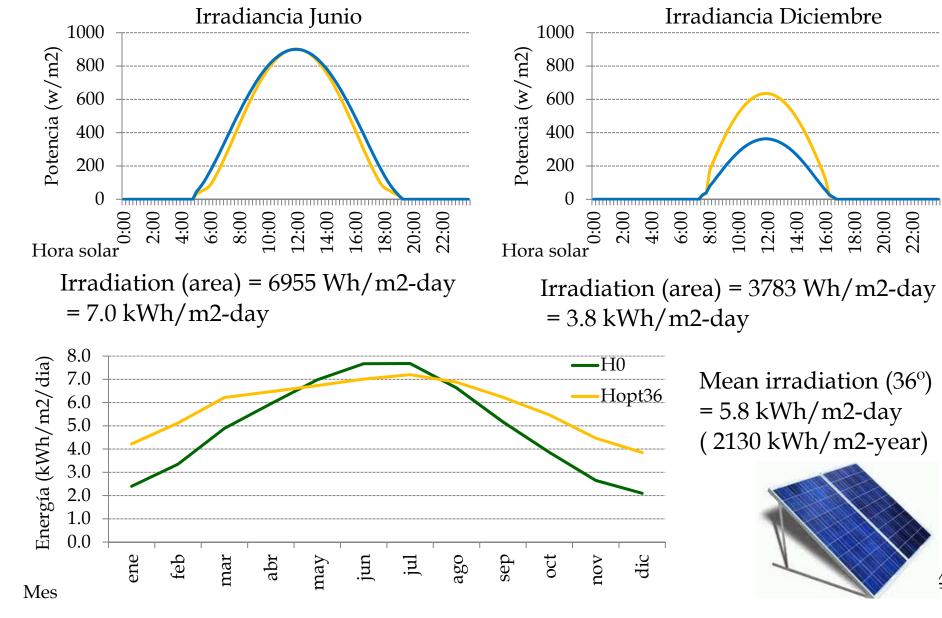


Irradiance (Power)-Irradiation (Energy) Valencia: fixed-horizontal (PV-GIS)

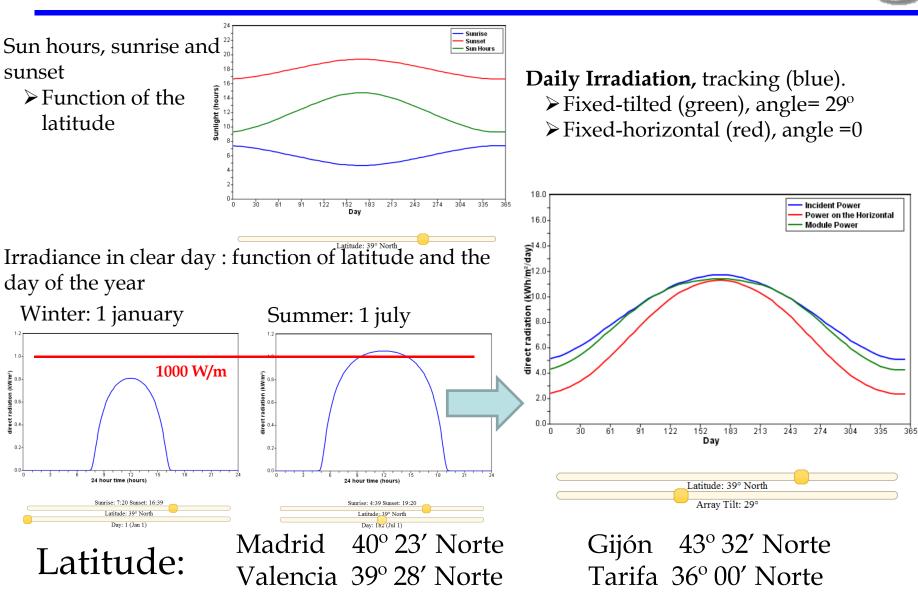




Irradiance (Power)-Irradiation (Energy) Valencia: tilted surface 36° opt (PV-GIS)



Irradiance (Power)-Irradiation (Energy)

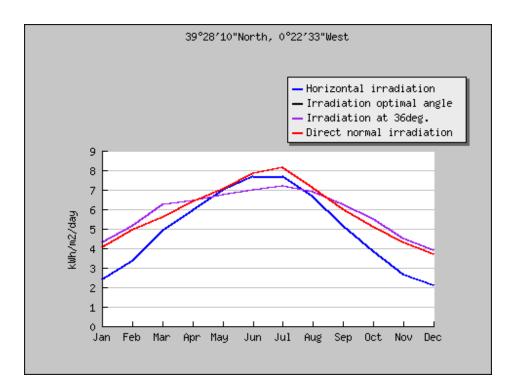


http://pveducation.org/pvcdrom/properties-of-sunlight/calculation-of-solar-insolation

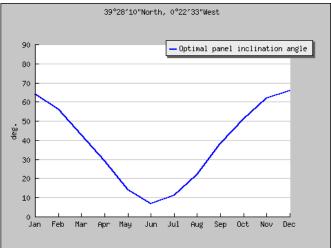
Solar Energy in Europa-PVGIS

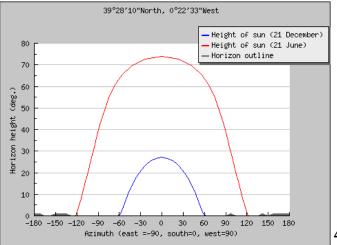


PV-GIS (Valencia, Latitude 39.5°): Horizontal: 4.96 kWh/m2/day – 1,810 kWh/m2/year Optima (36°): 5.85 kWh/m2/day - 2,135 kWh/m2/year



http://re.jrc.ec.europa.eu/pvgis/apps4/pvest.php#



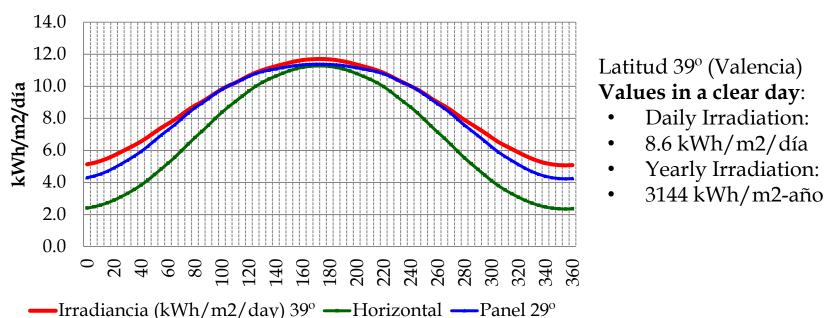


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Daily/Yearly Irradiation



Irradiance=> Power: energy that reaches in each instant Irradiation=> Energy: volume of energy accumulated in a day or in a year

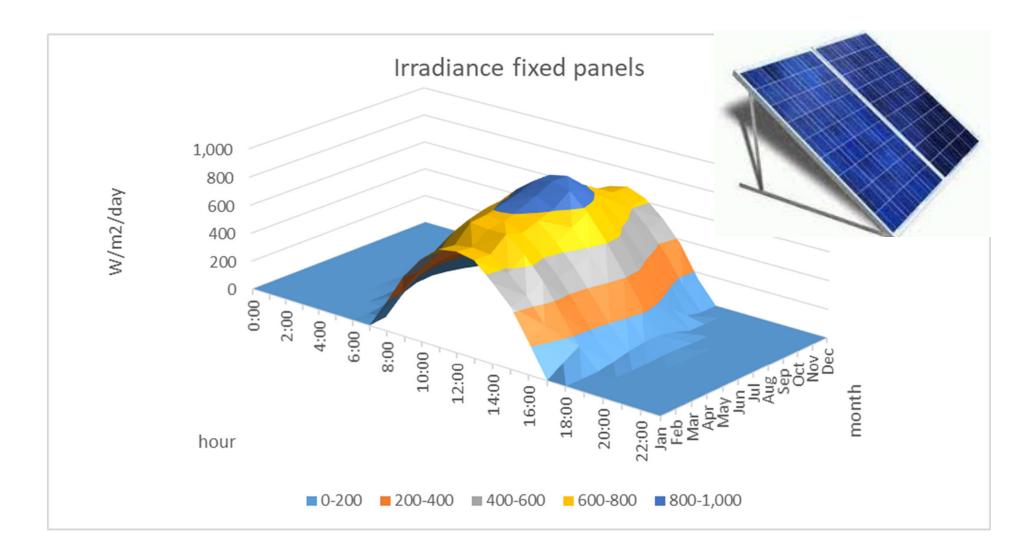


Irradiation

	Incoming Lat 39°	Horizontal	Tilted 29°
Daily Irradiation kWh/m2/day	8.6	6.8	8.2
Annual Irradiation kWh/m2-year	3,144	2,493	3,011
efficiency	0.75	0.75	0.75
Potential energy kWh/m2-year	2,358	1,870	2,258

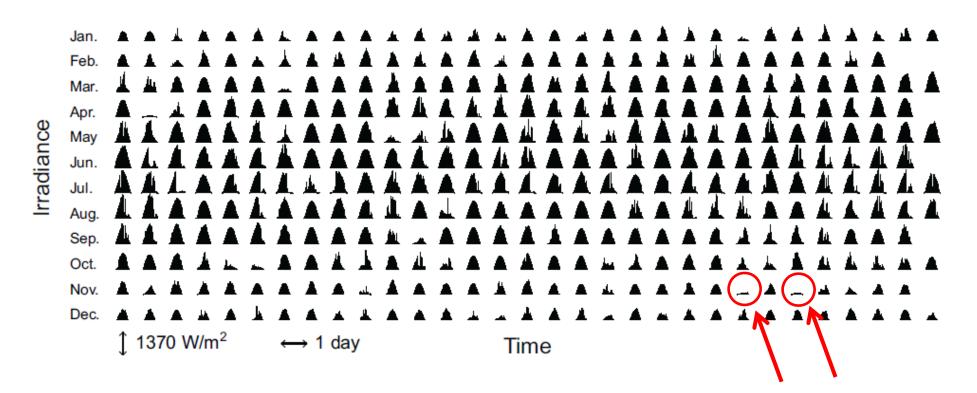
Irradiance optimal angle







Solar energy throughout the day for a full year There are days with almost zero irradiance all day. Need of connection to the network or the use of accumulators



Complete Solar Irradiance Profi le in Golden, Colorado for the Year 2012